

# Exploring Baroclinic Mode 2D ADCIRC to Capture Inter/Intra-annual Sea Surface Variations

Overall Concept, Implementation Using File I/O, and  
Preliminary Results

William Pringle and Joannes Westerink

Computational Hydraulics Laboratory, University of Notre Dame

Thursday, April 12, 2018

# Outline

Introduction

Method Description

Domain

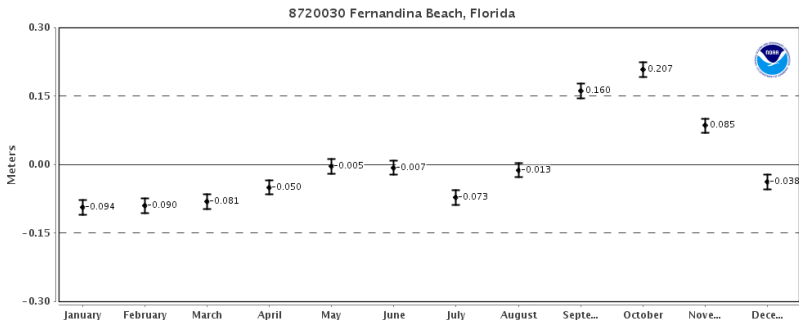
Dissipation due to Internal Tides

Baroclinic Pressure Gradient Terms

Fully Forced Yearlong Simulation

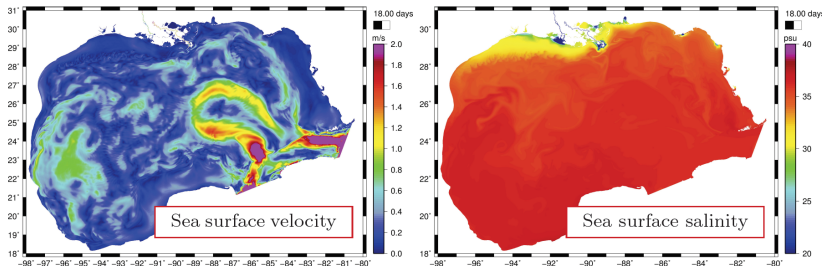
# Background

- ▶ Flooding risks for coastal cities change based on inter-annual and intra-annual variability
  - ▶ Long-term warming of the ocean
  - ▶ Seasonal warming and cooling
  - ▶ Changes in the Gulf Stream fluxes affect the sea surface height on the US East and Gulf Coasts
  - ▶ The interaction of winds and nearshore stratification also affect US East Coast water levels



# Problem

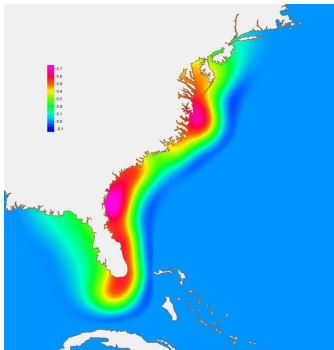
- ▶ The processes mentioned are primarily **baroclinic**, but ADCIRC storm surge analysis almost always computed in 2DDI barotropic mode
- ▶ 3D baroclinic ADCIRC has recently made great progress, but still difficult and horizontal resolution must be sacrificed



Source: A Fathi, J. C. Dietrich, C. N. Dawson, K. M. Dresback, A. Samii, R. Cyriac, C. A. Blain, R. Kolar. Prediction of surface oil transport in the Northern Gulf of Mexico by using a three-dimensional high-resolution unstructured-grid baroclinic circulation model. *Ocean Circulation*, 2017. *In Preparation*.

# Solutions

- ▶ Hindcast mode: geoid offset (constant or spatially varying) that can be tuned
- ▶ Forecast mode: pseudo atmospheric pressure field that drives an inverted barometer effect (Rick's talk last year)



- ▶ **This Study:** 2DDI baroclinic mode ADCIRC that is fed information from a widely used and validated operational 3D baroclinic model (e.g. HYCOM)

## Background: 3D Baroclinic Mode ADCIRC

- ▶ Splits-up the external (fast, barotropic) and internal (slow, baroclinic) modes because of the disparate timescales
- ▶ The 3D and baroclinic information is passed from the internal model to solve the external mode solution (and external mode also passed back to internal mode)
- ▶ Different time step for each mode can be used

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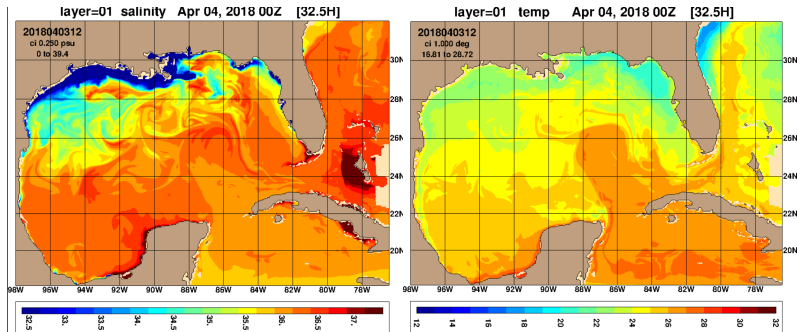
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- ▶ Different time step for each mode can be used

BUT, both modes are computed on the same horizontal grid...

**Difficult to obtain high horizontal resolution because of vertical dimension and many more computational operations than 2DDI**

# This Study: 2DDI Baroclinic Mode ADCIRC

- ▶ Do the mode splitting between **separate models** (e.g. HYCOM and ADCIRC)
- ▶ Allows us to obtain really high horizontal resolution at the coast (in the ADCIRC model) while absorbing the 3D baroclinic information
- ▶ Leverages on the quality of existing widely used and validated 3D baroclinic models





# Depth-integrated Momentum Equation

$$\begin{aligned} \frac{\partial \mathbf{u}}{\partial t} + (\mathbf{u} \cdot \nabla) \mathbf{u} + f \mathbf{k} \times \mathbf{u} = & -\nabla \left[ \frac{p_s}{\rho_0} + g(\zeta - \zeta_{EQ} - \zeta_{SAL}) \right] \\ & + \frac{\nabla M}{H} - \frac{\nabla D}{H} - \frac{\nabla B}{H} + \frac{\tau_s}{\rho_0 H} - \frac{\tau_b}{\rho_0 H} - \mathcal{F}_{IT} \end{aligned} \quad (1)$$

- Baroclinic pressure gradient:

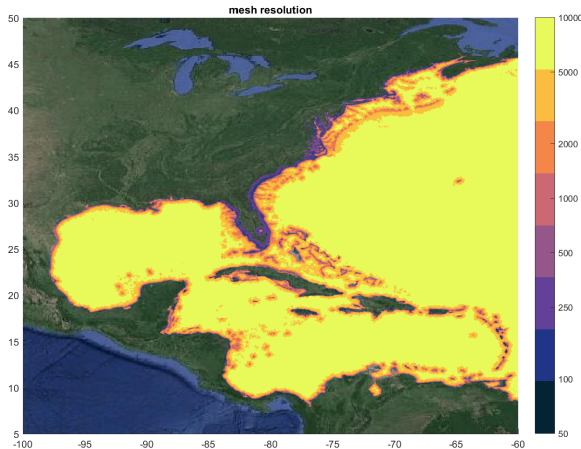
$$\nabla B = \int_{-h}^{\zeta} \left( g \nabla \left[ \int_z^{\zeta} \frac{\rho - \rho_0}{\rho_0} dz \right] dz \right) dz$$

- Internal tide induced barotropic energy dissipation:

$$\mathcal{F}_{IT} = C_{IT} \frac{[(N_b^2 - \omega^2)(\tilde{N}^2 - \omega^2)]^{1/2}}{\omega} (\nabla h \cdot \mathbf{u}) \nabla h$$

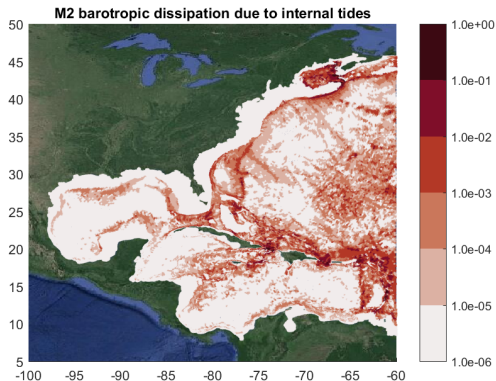
# Application: US East Coast Gulf Coast (ECGC)

- ▶ 2.46 million nodes
- ▶ ~60 m minimum element size at the coast
- ▶ Floodplain up to 10 m above LMSL
- ▶ Stable time step of **5 sec** in explicit model (full advective terms everywhere)



- ▶ Built using OceanMesh2D toolbox by Keith
- ▶ Come see us on Friday afternoon for hands-on session!

# Dissipation due to Internal Tides

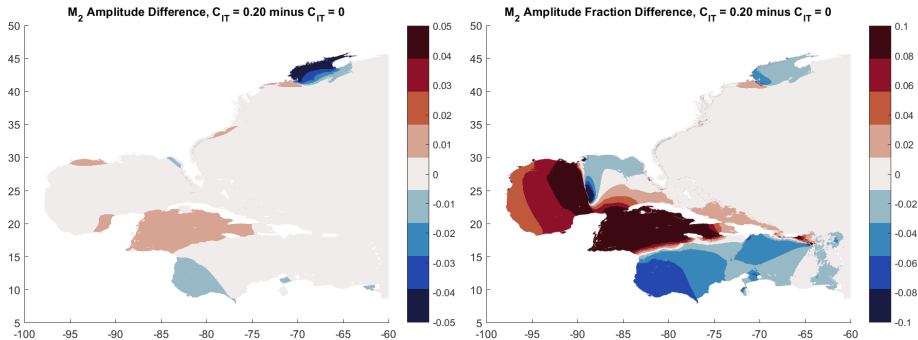


- ▶ Globally  $\approx 30\%$  of barotropic dissipation due to internal tide energy conversion
- ▶ ECGC domain  $\approx 14\%$  (11 GW out of 75 GW total)
- ▶ Gulf of Maine bed stress dominates
- ▶ No significant submarine ridges in domain

$$\mathcal{F}_{IT} = C_{IT} \frac{[(N_b^2 - \omega^2)(\tilde{N}^2 - \omega^2)]^{1/2}}{\omega} (\nabla h \cdot \mathbf{u}) \nabla h \quad (2)$$

$C_{IT}$ : tunable constant:  $\sim 0.2$ ,  $\tilde{N}$ : depth-averaged Brunt-Vaisala buoyancy frequency,  $N_b$ : Brunt-Vaisala buoyancy frequency at seabed,  $\omega$ : frequency of dominant tide ( $M_2$ ),  $h$ : ocean depth at rest,  $\mathbf{u}$ : depth-averaged velocity

# Effect of Internal Tide Dissipation on $M_2$ Tide



- ▶ Decreases amplitudes in Gulf of Maine and Southern Caribbean Sea
- ▶ Amplitudes actually increase in Gulf of Mexico and Northern Caribbean Sea

## Change in Integrated Average Error

$$D = \left( 0.5 \sum_k \left[ (A_0^k)^2 + (A_m^k)^2 - 2A_0^k A_m^k \cos(\theta_0^k - \theta_m^k) \right] \right)^{1/2}$$

$$\bar{D}_{tpx} = \frac{\iint D \, dA}{\iint dA}$$

**Table 1:** Integrated RMS Discrepancy versus TPX08,  $\bar{D}_{tpx}$

Region	Tidal Wave	Model	
		$C_{IT} = 0$ $\bar{D}_{tpx} (SD_{tpx})$	$C_{IT} = 0.20$ $\bar{D}_{tpx} (SD_{tpx})$
Deep $h > 500$ m	M <sub>2</sub>	1.25 (4.51)	0.74 (3.01)
	K <sub>1</sub>	0.27 (0.34)	0.21 (0.32)
	All	1.51 (4.52)	0.97 (3.06)
Shelf $h < 500$ m	M <sub>2</sub>	3.72 (7.32)	3.34 (7.26)
	K <sub>1</sub>	0.86 (1.35)	0.87 (1.36)
	All	4.74 (7.60)	4.37 (7.59)

# Baroclinic Pressure Gradient Term

- Fully depth-integrated term:

$$\nabla B = \int_{-h}^{\zeta} \left( g \nabla \left[ \int_z^{\zeta} \frac{\rho - \rho_0}{\rho_0} \right] dz \right) dz$$

- In ADCIRC User Guide, for 2DDI ADCIRC if  $\rho$  is assumed to be uniform over the depth (this formula is in timestep.F)...

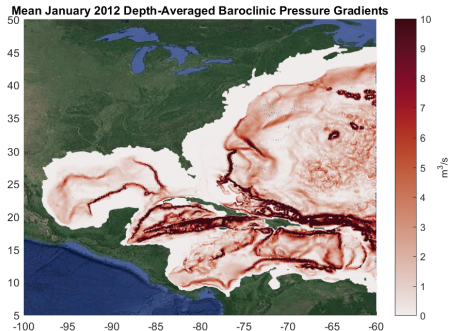
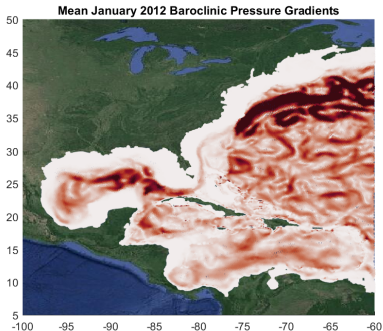
Depth-averaged term:

$$\nabla B = gH \left[ \frac{\rho_{2D} - \rho_0}{\rho_0} \nabla \zeta + \frac{H}{2} \nabla \left( \frac{\rho_{2D} - \rho_0}{\rho_0} \right) \right]$$

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- BUT,  $\rho$  is *not* uniform over the depth, leads to incorrect specification:



## Current Implementation

- ▶ The mode splitting is one-way (HYCOM  $\rightarrow$  ADCIRC)
- ▶ Compute baroclinic pressure gradient,  $\nabla B$  and internal tide dissipation parameter,  $\mathcal{F}_{IT}$  offline (I do this in MATLAB)
- ▶ Pass the various parameters through file I/O basically using existing ADCIRC options
  1. Made a IDEN = -5 option specified in **fort.15** and enabled a switch for 2DDI baroclinic mode



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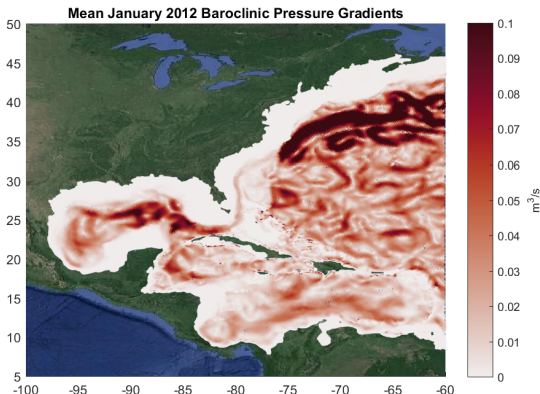
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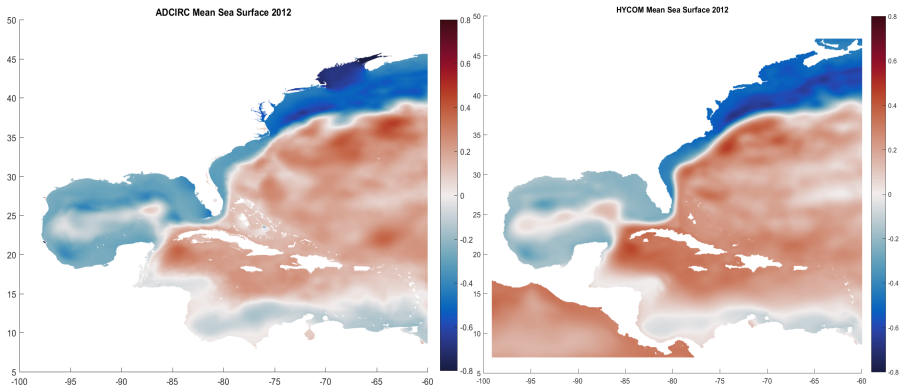
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  4. Open boundary elevation specified conditions are the periodic tidal frequencies (TPXO9 harmonic constituents) + non-periodic elevations specified from a **fort.19** (elevations are extracted from HYCOM, they do not simulate ocean tides)

# Fully Forced Yearlong Simulation (2012)

- ▶ Daily *HYCOM* + *NCODA Global 1/12° Reanalysis* elevations specified at the open boundary (added to tidal constituents)
- ▶ Daily baroclinic gradients calculated offline from 3D *HYCOM* temperature and salinity fields
- ▶ Hourly  $0.281^\circ$  ERA5 atmospheric data
  - ▶ Spin-up plus relaxation period of 30 days



# Yearly Mean Sea Surface Height

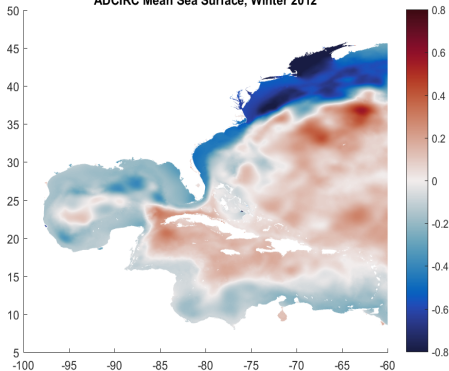


- ▶ General ocean surface patterns are reproduced
- ▶ Daily outputs but ADCIRC includes tides so average may be slightly biased

# Seasonal Mean Sea Surface Height

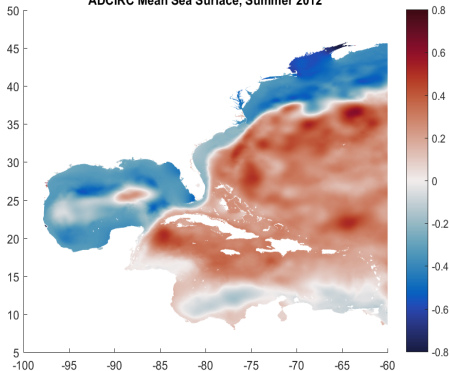
Winter (Nov - March)

ADCIRC Mean Sea Surface, Winter 2012



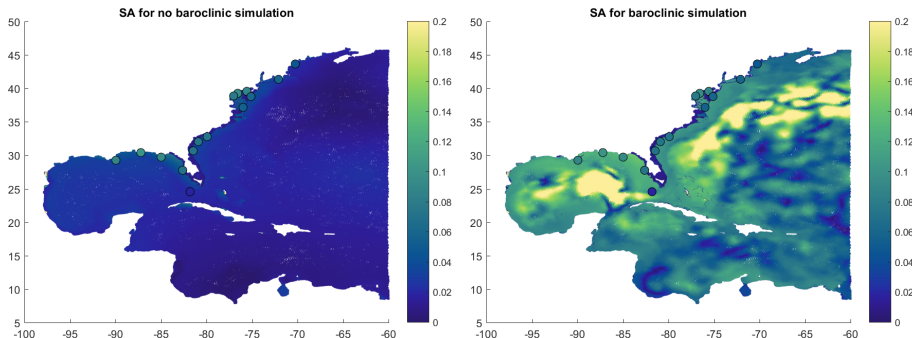
Summer (June - Oct)

ADCIRC Mean Sea Surface, Summer 2012



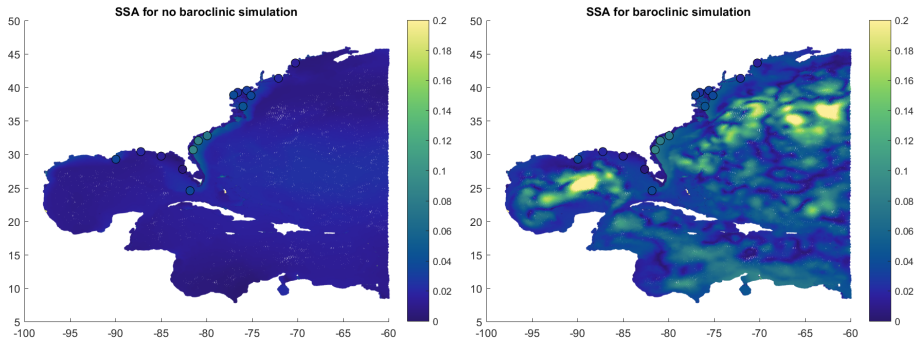
- Able to get higher sea surface in summer than winter in general

# SA Constituent Map



- ▶ No baroclinic mode induces notable SA only in the Florida panhandle coastal region (direct seasonality of winds and pressure)
- ▶ Baroclinic mode induces much larger SA in general due to the changes in the Gulf Stream, heating-cooling, and warm core eddies

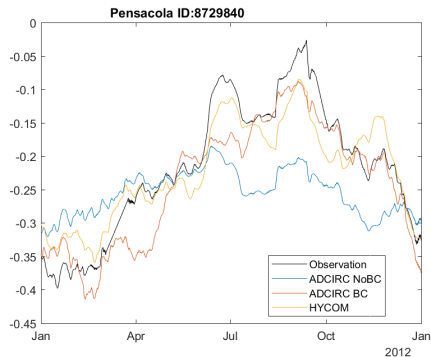
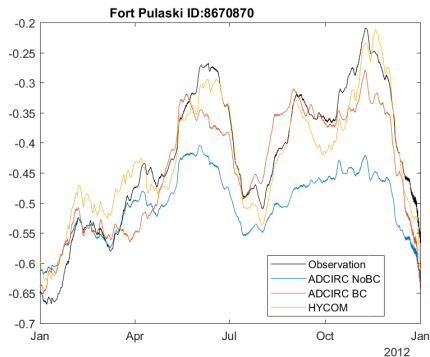
# SSA Constituent Map



- ▶ No baroclinic mode induces large SSA only in the South Atlantic bight
- ▶ Baroclinic mode larger SSA, particularly in the Gulf and the northern Atlantic

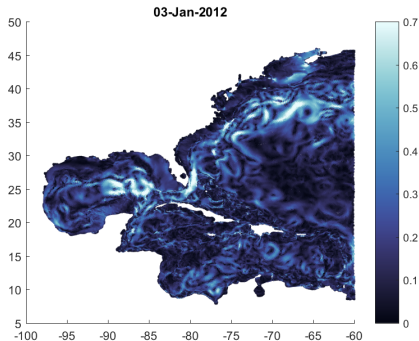
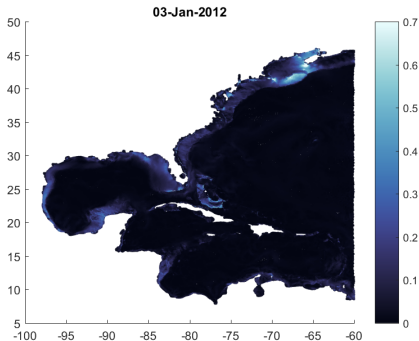


# Yearly 30-day Moving Mean Time Series at Tide Gauges



- ▶ From July onwards baroclinic mode can capture the seasonal rise that non baroclinic model cannot
- ▶ Around April the baroclinic mode misses something, problems:
  - 1) Open ocean boundary (mis-match between ADCIRC & HYCOM)?
  - 2) Nearshore/estuarine density gradients?

# Ocean Currents



- ▶ Geostrophic ocean currents are generated by baroclinic pressure gradient terms that are otherwise not present in non-baroclinic 2DDI ADCIRC

# Conclusions

- ▶ Representing around 14% of  $M_2$  tidal dissipation, internal tide induced dissipation improves tidal solutions in ECGC region
- ▶ Inclusion of baroclinic pressure gradient terms in 2DDI ADCIRC drive geostrophic ocean currents
- ▶ Geostrophic ocean currents in 2DDI ADCIRC induce mean sea surface heights similar to 3D baroclinic HYCOM model
- ▶ Greater intra-annual variability in baroclinic mode versus non-baroclinic mode demonstrated (e.g. SA, SSA)
- ▶ Complete hydrodynamic modeling system that does not require ad-hoc data manipulation

# Going Forward

Problems to solve:

- ▶ May need to reconcile discrepancy between ADCIRC and HYCOM at open boundary through e.g. sponge type layer
- ▶ What to do with densities in nearshore/estuarine areas where HYCOM has no data/resolution?

Online implementation:

- ▶ Directly read in HYCOM temperature and salinity to calculate  $\nabla B$  and  $\mathcal{F}_{IT}$  online (in ADCIRC)
- ▶ Make use of ESMF framework and libraries for interpolation
- ▶ **3D** terms?? (dispersion, mixing, bottom stress)
- ▶ Final aim is ESMF compliant operational model, possibly **two-way coupled** with a 3D baroclinic model (ADCIRC can be used to provide 3D model with accurate coastal elevations)